

# Exploring the use of Unmanned Aerial Vehicles (UAVs) with the simplified “triangle” technique for Soil Water Content and Evaporative Fraction retrievals in a Mediterranean setting

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## ABSTRACT

Information acquired from Unmanned Aerial Vehicles (UAVs) is frequently used nowadays in a variety of disciplines and research fields. The present study explores for the first time the combined use of UAVs with a newly proposed technique for estimating evaporative fraction (EF) and surface soil moisture (SSM). The investigation is performed in a typical Mediterranean setting, a citrus field with flat topography divided in two plots with different irrigation schemes, in Sicily, Italy, at which ground data acquired during an extensive field campaign in July 2019. Reasonable estimates of both EF and surface wetness were produced, with patterns in agreement to vegetation cover fragmentation, topography, and other site-specific characteristics. Validation shows average error of 0.053 for EF and of 0.040 cm<sup>3</sup> cm<sup>-3</sup> for SSM. The results are comparable or better to those reported in analogous studies performed in similar areas. This implies that the investigated approach performs well under the semi-arid conditions characterising the experimental set up. To our knowledge, this study represents the first evaluation of the combined use of the “simplified triangle” with very high-resolution UAV imagery. As such, the findings are of significance regarding the potential future use of the “simplified triangle” approach particularly with very fine resolution imagery such as that provided by UAV for mapping and monitoring EF and SSM in agricultural and natural ecosystems.

**KEYWORDS:** earth observation, unmanned aerial vehicles, surface soil moisture, evaporative fraction, simplified triangle, surface temperature/vegetation index

## 1. Introduction

The natural processes taking place on the Earth’s surface control the energy and mass exchanges between land and atmosphere and are key drivers of the Earth’s system (North et

al., 2015; Gerken et al., 2019). Today, particularly so in light of climate change and concerns related global food and water security, an improved understanding of land-atmosphere interactions is a topic of urgent importance (Ireland et al., 2015; Deng et al., 2019). In this context, obtaining accurate information on the spatial and temporal variability of land surface parameters such as evaporative fraction, EF (defined as the ratio of instantaneous latent heat flux (LE) to net radiation ( $R_n$ ) and surface soil moisture (SSM) is of primary interest for several environmental applications and research investigations (Jung et al., 2011; Srivastava et al., 2019). This is due to the influence of these parameters on key physical processes and feedback loops of the Earth system (Nutini et al., 2104; Srivastava et al., 2015; Amani et al., 2016). Accurate information on their spatiotemporal variability, particularly at fine spatial and temporal resolution, can provide valuable information in research studies and practical applications linked to ecosystem processes, plant water requirements and water resources management (Shi et al., 2014; Minacapilli et al., 2015; Deng et al., 2019; Yang et al., 2020).

Despite their significance, it is quite difficult to quantify EF and SSM on a routine basis over large geographical regions using ground instrumentation. The main reasons include the large spatiotemporal variability of these parameters (Bao et al., 2018). Earth Observation (EO) presents a suitable alternative to ground observations for deriving SSM and/or EF over large regions and diverse geographical scales (Tian et al., 2014). A variety of approaches have been proposed for this purpose, ranging from semi-empirical to physically-based ones (see Petropoulos et al., 2015; 2018). Those approaches are characterised by different degree of complexity, input parameters requirements and retrieval accuracy.

A specific group of EO-based techniques commonly termed as surface temperature ( $T_s$ ) and vegetation index (VI) methods ( $T_s$ /VI), has shown an excellent promise at deriving spatially explicit maps of sensible and latent heat fluxes (H, LE) and/or SSM. These methods utilise optical (visible and infrared - VNIR) and thermal infrared (TIR) EO data and are based on physical relationships between the satellite-derived  $T_s$  and a VI, the latter being associated to the existent degree of vegetation (Zhang et al., 2014; Capodici et al., 2020). If these parameters are in a scatter plot, provided that there is a full variability in VI, a triangular/trapezoidal shape similar to that shown in **Figure 1** emerges. This shape, characterised by the physical boundaries also shown in **Figure 1**, results from the  $T_s$  sensitivity to water content, which increases as a function of the proportion of bare soil exposed. The biophysical properties included in this  $T_s$ /VI domain are well-documented (Gillies et al., 1997; Chauhan et al., 2003; Maltese et al., 2015; Wang et al., 2018; Cui et al., 2020). Detailed descriptions of these properties, including the key parameters affecting the  $T_s$ /VI scatterplot shape, are summarised in Petropoulos et al. (2009) and Petropoulos et al. (2018). Tang et al. (2017) introduced the End-member-based Soil and Vegetation Energy Partitioning model (ESVEP), a two-source approach for estimating land surface evapotranspiration (ET) ) for which two dry edges could be considered in the case of a root zone water stress occurs. It is based on the consideration that soil evaporation primarily draws water from the upper soil layer, whereas, transpiration exploits water from the root zone. The temporal response of soil water content of the upper soil and root zone in the framework of the ET process is therefore different: the dynamic of the soil water content is more rapid in the upper layer; it is slower in the root zone.

Recently, Carlson & Petropoulos (2019) proposed a  $T_s$ /VI technique for estimating both EF and SSM, which they named “simplified triangle”. This approach is essentially a variant of the so-called “triangle” technique (Carlson, 2007) and does not require for its implementation a

land biosphere model or any other ancillary data. These characteristics make this approach easy to apply in comparison to other  $T_s/VI$  methods. Furthermore, being dependent on a small number of easily computed EO-based parameters, it becomes a very attractive choice for potential operational use. Fuzzo et al. (2019) demonstrated how this newly introduced approach can be coupled with a crop prediction and a climatological water balance model in soybean yield prediction using MODIS data. However, as the technique is recent, studies validating its performance in different environments and with a range of EO instruments are scarce.

[Please put Figure 1 around here]

To our knowledge, this newly proposed technique has not been implemented on and verified for unmanned aerial vehicles (UAVs) data yet. UAV platforms with on-board visible/near infrared and thermal sensors have very important advantages over satellite EO platforms, such as user flexibility to select the target area and the frequency of data acquisition (Dawson et al., 2019; Liu et al., 2020). Therefore, this technique implementation with UAVs would be indisputably of key importance, as it would inform on its potential usefulness in a broad spectrum of practical applications and research purposes alike.

In this context, this study aims at exploring for the first time the combined use of the “simplified triangle” with very high spatial resolution UAV data, to predict the spatio-temporal variability of both EF and SSM. For this purpose UAV, ground truthing and ancillary data acquired during a field campaign that took place in July 2019 at one experimental site in Sicily, Italy, are employed. The experimental set up description is provided in Section 2, whereas the “simplified” technique implementation with the UAV data is made available in Section 3, followed by the results and the related discussion which are described in Sections 4, and 5, respectively.

## 2 Materials

### 2.1 Study site

The study site is a citrus orchard field (*C. reticulata* Blanco, cv. Tardivo di Ciaculli) located in the neighbourhood of Palermo, Italy (38° 4'53.4"N, 13° 25' 8.2"E). The site contains 30 year old tangerine trees planted at a regular spacing of 5.0 m × 5.0 m (plant density of 400 plants per ha) and irrigated with a subsurface drip system. The area is in a typical eastern Mediterranean semi-arid environment. The study area has flat topography with elevation between 30 and 35 m above sea level, and slopes ranging from 1% to 4%.

To differentiate irrigation management, the field has been divided into two plots of about 4,000 m<sup>2</sup> each, as shown in **Figure 2**. The first plot was maintained under full irrigation (FI), whereas the second under deficit irrigation (DI) applied throughout phase II of fruit growth (from 1 July 2019 to 20 August 2019). Each plot was, in turn, divided into four sub-plots differentiated for the anti-root agents introduced into the emitters during the manufacturing process, but not for the irrigation management, nor for the emitters' hydraulic performance. The subsurface drip system is characterized by two lateral pipes per plant row, installed at a distance of 1.1 m from the trees and buried at a depth of 0.30 m. In each lateral pipe, self-compensating emitters

were installed with half-meter spacing between them, nominal flow rate of  $2.3 \text{ h}^{-1}$  and operating pressure of 150 kPa. A disc filter, an electric control valve, a relief valve, a pressure gauge, and a flow meter completed each sub-plot irrigation unit.

The experimental setup is equipped with a WatchDog 2000 weather station (Spectrum Technologies, Inc.), including sensors for relative air humidity, wind speed and direction, air temperature, solar radiation, and rainfall, as well as eight "drill & drop" frequency domain reflectometry sensors (Sentek Pty Ltd, Stepney, Australia) to monitor soil water content, installed on a central tree of each sub-plot, 0.30 m away from the closest emitter. All the sensors were interfaced with a communications board that uses the cellular 3G data network for internet connection using the MODBUS RTU protocol to transfer and save the data into a MySQL database operated by AgriNET/Tuctronics which is accessible from the web. The system allows the download of weather variables, soil water content (SWC) and temperatures (T) in the root zone, at 10 cm intervals from the first 5 cm of the soil layer down to a depth of 0.6 or 1.2 m. The Scholander chamber (Scholander et al., 1965) was used to follow the temporal dynamic of predawn and midday stem water potential, whereas a couple of Granier thermal dissipation probes (Granier, 1985) was installed in four trees to monitor sap flow during the irrigation season.

In addition, an eddy covariance flux tower was set up in the orchard in February 2019 to measure the turbulent fluxes (sensible, H, and latent, LE, heat fluxes) and a four-component net radiometer was used to measure net radiation ( $R_n$ ) individual components. A CNR1 four component Net Radiometer was installed at 3.1 m a.g.l, while an InfraRed Gas Analyzer IRGA LI7500 (manufactured by LI-COR, Inc.) and a CSAT3 Three Dimensional Sonic Anemometer anemometer (manufactured by Kipp & Zonen B.V.) were installed slightly above, at 3.5 m above ground level (a.g.l.), , i.e., approximately 55 and 95 cm above the vegetation canopy. All the data were processed at 30 minutes interval. The footprint flux tower was calculated according to Schuepp et al. (1990) at 70% of the fluxes.

[Please put Figure 2 around here]

## 2.2 Data Acquisition & Pre-processing

### 2.2.1 Data Acquisition

The fieldwork for this study was carried out on July 2019. A series of spatial and ancillary data was acquired on 30 July 2019 as part of the field campaign that was conducted in order to support the study implementation. In particular:

- **Global Navigation Satellite System (GNSS) Survey.** Nine black and white control targets, and the same number of aluminium targets were distributed on a regular grid to cover the whole study area.

The coordinates of the targets were measured by a NRTK survey using a Topcon Hiper V receiver (both Global Positioning System (GPS) and Glonass constellations). A UNIPA (University of Palermo) GNSS Cross-origin resource sharing (CORS) network encompassing 8 permanent stations, 2 of them installed on two University buildings in Palermo and Agrigento and 6 at other public institutions of the Sicilian territory was employed for Network real-time kinematic (NRTK) positioning. The network covers about  $7400 \text{ km}^2$  western Sicily. The GNSS CORS Network project was carried out with the technical collaboration of Topcon Italy (that supported the scientific

research with GNSS receivers and antennae), in the framework of developing a network for technical (real-time) and scientific (post-processing) use. The CORSs included in the Topcon Netgeo GNSS network. Since 2013 the data retrieved from UNIPA GNSS CORS network have been used for the computation of the RDN2 (*Rete Dinamica Nazionale 2*) which provides the WGS84 datum for Italy in the European Permanent Network (EPN subnetwork). UNIPA GNSS CORS network has received the scientific acknowledgment through many experiments in various application fields (Catania et al. 2020; Angrisano et al. 2020, Kenyeres *et al* 2019, Pipitone et al. 2018, Dardanelli et al. 2015, Dardanelli et al. 2014, and Dardanelli and Carella, 2013). Since 2013 the postprocessing RINEX (Receiver INdependent EXchange) data have been made available for the evaluation of the national reference framework by the IGMI (Italian cartographic military institute) and for technical researches able to investigate the horizontal and vertical velocity map in Italy (Maseroli, 2015). NRTK positioning was carried out using the hardware and software infrastructure of the permanent Netgeo-Topcon Italy network framed in the reference system ETRF2000 (powered by UNIPA GNSS CORS) and in particular via the VRS (Virtual Reference Station) stream. Data availability and geodetic framework are described in Dardanelli *et al.* (2020). The processing of GNSS data acquired to allow an accurate orthorectification of multispectral and thermal images was carried out by Meridiana software ver. 2020.

- **Proximity sensing images.** Multispectral images were acquired using a NT8 contras octocopter carrying a RikolaDT-17 Fabry-Pérot camera (manufactured by Rikola Ltd). The multispectral camera has a 36.5° Field of View. It was set-up to acquire images in 9 spectral bands with a 10 nm bandwidth. Central wavelengths were 460.43, 480.01, 545.28, 640.45, 660.21, 700.40, 725.09, 749.51 and 795.53 nm. At a flight altitude of 50m above ground (a.g.l.), the average Ground Sampling Distance (GSD) was 3 cm. Thermal images were acquired almost simultaneously to the multispectral images, using a DJI Mavic 2 Enterprise Dual quadcopter carrying on-board a FLIR Lepton® (manufactured by FLIR® Systems, Inc) acquiring in the longwave infrared spectral range (from 8 to 14 µm), with a thermal sensitivity lower than 50 mK (0.050 °C). The average GSD was 3.46 cm. All the images were resampled at 4 cm spatial resolution using a pixel aggregate resampling method.
- **Spectroradiometric measurements.** Four reference targets, ranging in a greyscale from black to white were also positioned to allow the spectral reflectance calibration by means of a field spectroradiometer. The employed ASD FieldSpec®FR spectroradiometer (Analytical Spectral Device, ASD, Inc.) measured the full solar spectrum (between 350 and 2500 nm) with no fore optic attached.
- **Thermographs.** Ground measurements of surface temperature ( $T_s$ ) were carried out at noon using a handheld FLIR SC660 (FLIR® Systems, Inc.) characterized by a sensitivity lower than 30 mK.

## 2.2.1 Pre-processing

Following the data acquisition, standard pre-processing steps were applied. To orthorectify the multispectral and thermal images, a standard photogrammetric/SfM approach (e.g., Harwin and Lucieer, 2012) was applied via Pix4D mapper (by Pix4D Inc.). A Topcon Hiper V receiver (both GPS and GNSS Connectivity) was employed to acquire ground control points for the orthorectification. The average position dilution of precision (PDOP) and the geometric dilution of precision (GDOP) were 1.8 and 2.0, respectively. The control targets were positioned with

average planimetric and altimetric accuracy of  $\pm 2$  cm that can be considered within acceptable geometrical configuration limits to orthorectify the UAV images, considering that these latter are characterized by a spatial resolution of 4 cm once orthorectified. Images acquired in the visible and near infrared were calibrated to ground reflectance implementing the empirical line technique (Karpouzli and Malthus, 2003), which allows the simultaneous correction of the atmospheric influence. Similarly, TIR images were calibrated into surface radiometric temperatures by means of a linear regression with at ground thermographs and an emissivity map of the soil vegetation system (Negm et al., 2017). The spatial distribution of emissivity was calculated according to Valor and Caselles (1996). Given the spatial resolution of the images (about  $10^{-2}$  m) compared to the spacing of the trees (about 5 m) we did not consider the cavity effect. We assume the emissivity values for bare soil and densely vegetated ground to be equal to 0.97 and 0.99, respectively, as reported in Sobrino et al. (2004). **Figure 3** illustrates the Normalized Difference Vegetation Index (NDVI) and of Surface Temperature ( $T_s$ ) final products upon completion of all pre-processing steps.

[Please put figure 3 around here]

### 3 Methods

#### 3.1 Simplified Triangle Method

A comprehensive account of the “simplified” triangle technique implementation is available in Carlson and Petropoulos (2019). Briefly, the method allows the retrievals of two parameters, the soil water availability ( $M_o$ ) and EF.  $M_o$  represents surface wetness in the bare soil surface (top few millimetres of it) and it is computed from the ratio between the actual soil/vegetation system evapotranspiration ET and potential evapotranspiration ( $ET/ET_p$ ).  $M_o$  is also equated to SSM by multiplying  $M_o$  with the soil’s field capacity. On the other hand, EF is defined as the ratio between latent heat flux (LE) and net radiation ( $R_n$ ).

EF and  $M_o$  are obtained from the  $T_s/VI$  feature space. The scatterplot is constructed by plotting the  $T_s$  versus fractional vegetation ( $F_r$ ), where the latter is computed from the NDVI (see Equation (1) below) and its corresponding range of variability, as proposed by Carlson (2007). Upon completion of this step, a number of parameters need to be determined, namely: (a) the NDVI values for bare soil and dense vegetation (respectively,  $NDVI_o$  and  $NDVI_s$ ), and (b) the highest value of  $T_s$  ( $T_s [max]$ ) which is characteristic of dry/bare soil pixels, as well as the minimum value of  $T_s$  ( $T_s [min]$ ).

$NDVI_o$ ,  $NDVI_s$ ,  $T_{max}$  and  $T_{min}$  are used to specify the  $T_s/VI$  feature space boundaries and to constrain the solution for EF and  $M_o$ .  $NDVI_s$  and  $T_{min}$ , represent dense vegetation and define the lower left (wet) vertex of the triangle, i.e. the so-called ‘wet edge’ or ‘cold edge’ (see **Figure 4**). The wet edge corresponds to  $M_o$  and EF values equal to 1.0. Similarly,  $NDVI_o$  and  $T_{max}$  define the lower right vertex of the triangle, the so-called ‘dry edge’ or ‘warm edge’ (also shown also in **Figure 4**). These points characterize the soil dryness boundary with  $M_o = 0$  and covers the area from  $T_{max}$  and  $NDVI_o$  to  $NDVI_s$ , which, for a triangle with a distinct upper vertex, occurs at  $T_{min}$ . Even though  $M_o = 0$  along the “dry edge”, along the dry edge EF itself is non-zero apart from the triangle’s lower right vertex. The next step in the technique implementation includes the scaling of  $T_s$  to  $T^*$  (by applying Equation (2) below), which ranges between zero to one.

At this stage two central hypotheses are made. The first is that when vegetation is at wilting point transpiration is always equal to the potential transpiration, as generally assumed in nearly all  $T_s/VI$  approaches (e.g., Jiang and Islam 2003). The second hypothesis is related to the relationship between EF and  $M_o$  within the  $T_s/VI$  domain, which is assumed to be linear.

[Please put Figure 4 around here]

Thus, on the basis of the assumptions above,  $M_o$  is defined as the ratio between the lengths “a” and “d”. Both these lengths depend on  $T^*$  and  $F_r$ . For conditions where a pixel comprises of both areas of vegetation and bare soil, the canopy EF is taken as the weighted value of EF for the vegetation fraction of the pixel ( $EF_{veg} = 1$ , by definition). As such, both  $M_o$  and EF are computed for all pixels contained in the  $T^*/F_r$  domain from the implementation of Equations (3) and (4) shown below.

$$F_r = \left\{ \frac{(NDVI) - (NDVI_0)}{(NDVI_s) - (NDVI_0)} \right\}^2 \quad (1)$$

$$T^* = \{T - T_{min}\} / \{(T_{max} - T_{min})\} \quad (2)$$

$$M_o = 1 - T^*(pixel)/T^*(dry\ edge) \quad (3)$$

$$EF = (EF_{soil})(1 - F_r) + F_r (EF_{veg}) = M_o(1 - F_r) + F_r \quad (4)$$

In the above,  $EF_{soil}$  refers to the ratio between soil evaporation and net radiation.  $T^*(pixel)$  is the scaled surface temperature  $T^*$  for a given pixel within the scatterplot and  $T^*(dry\ edge)$  is the value of  $T^*$  at the dry edge of the triangle. In this study, the values for the temperature limits were  $T_{min} = 19.40\ ^\circ C$  and  $T_{max} = 73.27\ ^\circ C$ , whereas for NDVI were  $NDVI_0 = 0$  and  $NDVI_s = 1$ . Noticeably that fully vegetated pixels exhibit a variability in  $T^*$  of 0.25 conferring to the  $T^* - F_r$  scatterplot a trapezoidal shape. The variability in  $T^*$  could be attributed to the very high spatial resolution achieved by UAV which allows to record the surface temperatures of the single leaves of the same canopy. In particular, the variability in  $T^*$  is attributed to the different exposure to the direct solar radiation of the single leaves which controls i) directly, the individual leaf warming up; ii) indirectly, the leaf transpiration.

The implementation of the steps summarized above to the pre-processed UAV data resulted in the scatterplots of NDVI vs  $T_s$  and of computed  $F_r$  vs  $T^*$  shown in **Figure 5**. The spatial maps of  $F_r$  and  $T^*$  are also shown in this figure.

[Please put Figure 5 around here]

### 3.2 Statistical Analysis

Evaluation of the predicted SSM and EF included at first a visual inspection of the spatiotemporal variability of the derived maps. Next, the main validation approach involved comparisons at pixel level between the predicted and measured parameters. The statistical scores computed that quantify the agreement between predictions and observations are summarised in Table 1. These statistical measures have already been used in similar past verification exercises (e.g. Nutini et al., 2014; Piles et al., 2016; Amani et al., 2016; Xu et al., 2018, Wang et al., 2018).

[Please put Table 1 around here]

## 4 Results

### 4.1 Visual Comparisons

The EF and SSM maps and their corresponding histograms obtained from the UAV data and the “simplified triangle” technique are illustrated in **Figure 6**. The first step of the analysis included a visual inspection of the spatial variability of the derived parameters. As can be observed, both EF and SSM maps exhibited a sensible range of values as well as reasonable spatial variability. Clearly, the spatial variability is in agreement with the changes in land use/cover, as well as with the derived  $F_r$  and  $T_s$  maps based on the UAV data that were presented in **Figure 5**. Both EF and SSM predicted by the “simplified” triangle are spatially consistent with the soil/vegetation cover patterns and variability: in particular, high values of both variables correspond to the vegetated areas of the image, whereas low values appear in areas of bare soil.

To further illustrate the above observation, it was further investigated the variability of the derived parameters separately for the bare soil and the partially or fully vegetated components (see **Figure 7**). As evidenced in the maps shown in **Figure 7** (and their associated histograms), the variability of the examined parameters is largely explained by the spatial variability in the land surface fragmentation. It is evident from the visual comparisons of bare soil and vegetation maps and histograms, that the variability of the vegetation for both EF and SSM is significantly higher in bare soil. From these figures it is shown that the EF and SSM for vegetation are predominantly above 0.9 EF and 0.2 SSM. Bare soil presents higher variability, but the highest frequencies (especially for SSM) are close to 0.26.

[please put figure 6 around here]

[Please put figure 7 around here]

The last step of the visual analysis focused on an arbitrary transect, chosen as the diagonal line connecting the North and the South vertices of the experimental site. The spatial evolution of each predicted parameter along this transect is depicted in **Figure 8**. This approach allows examining simultaneously the variability of the different parameters, namely of EF, SSM,  $F_r$ , and  $T^*$ . The results of this analysis are depicted in **Figure 8**. As one can notice, the variability of the predicted parameters within the field follows largely explainable trends, depending on both  $F_r$  and  $T^*$ . This observation provided further evidence of the technique’s ability to satisfactorily predict both EF and SSM in the field when implemented with the UAV data.

[Please put figure 8 around here]

### 4.2 Point Comparisons

The results which concerned point-wise (i.e. pixel level) comparisons are summarised in **Table 2**. As already noted, ground measurements of the radiation and turbulent fluxes were acquired at a single location within the experimental field. On the other, SSM measurements were conducted at a total of eight sites across the field, in which two different irrigation



strategies were applied since 1 July 2019. In particular, sites 1 to 4 were maintained under full irrigation, whereas sites 5 to 8 under water deficit conditions.

As can be observed (in **Table 2**), the “simplified triangle” achieved very good predictions of both EF and SSM, which are in close agreement to the field observations and in the same range as the results of similar studies (e.g., Peng and Loew, 2014; Bai et al., 2019). The predicted EF value, compared with the observed one, was slightly overestimated, with an absolute difference of 0.053. However, it should be noted that this difference is also based on a single ground measurement, since there was only one eddy covariance station installed in the central part of the experimental site. In reference to the soil water content, Table 2 shows that the predicted SSM is in very good agreement with the respective measurements, with RMSE of 0.040 cm cm<sup>-3</sup>. Scatter (0.031 cm cm<sup>-3</sup>) contributes to RMSE relatively more than Bias (-0.025 cm cm<sup>-3</sup>) but not overly so.

[please put Table 2 around here]

As shown in **Table 2**, the mean predicted SSM (denoted as “P”) for the locations of Stations 1 to 4 (plots with full irrigation) is 0.123 cm cm<sup>-3</sup> while for locations of stations 5 to 8 (plots with deficit irrigation) the mean predicted SSM is lower at 0.096 cm cm<sup>-3</sup>. On the other hand, the measured SSM by the stations (denoted as “O”) does not reveal remarkable differences between plots maintained under different irrigation strategies. The mean observed SSM for plots 1 to 4 is 0.138 cm<sup>3</sup> cm<sup>-3</sup>, while for the plots 5 to 8 it is only marginally lower and equal to 0.131 cm cm<sup>-3</sup>. While bias is generally low, the predicted SSM underestimates the corresponding values in all the plots under deficit irrigation by -0.035 cm cm<sup>-3</sup> on average. However, for the fully irrigated plots, the underestimation is less than half in magnitude (equal to -0.015 cm cm<sup>-3</sup>). All in all, these results suggest that the “simplified triangle” performed satisfactorily in predicting both the EF and SSM under the examined conditions.

## 5. Discussion

Based on the results obtained (Section 4), the “simplified triangle” technique performed well to in reproducing the high spatial resolution of EF and  $M_o$ /SSM maps for the study area. Both predicted maps exhibited a largely explainable spatial variability across the experimental site, with patterns in agreement to land cover type, topography and other site-specific characteristics. In terms of statistical agreement, prediction accuracy was good for both EF and SSM, and in agreement to the accuracies reported by other independent investigators using different approaches and EO data types. For EF the difference between the predicted and measured value is 0.053, giving a slight overestimation. After the  $M_o$  was converted to SSM for the 8 stations, the results showed fairly low RMSE (0.040 cm cm<sup>-3</sup>) and low underestimation (Bias = -0.025). These values are close to those reported by other studies retrieving EF and SSM using TIR-based techniques (e.g., Peng and Loew, 2014; Nutini et al., 2014; Lu et al., 2015; Xu et al., 2018; Bai et al., 2019). Thus, findings, although are based on the single image analysis, are confirming the usefulness of the examined technique for EF and SSM spatial determination at very fine resolution when implemented with UAV data.

There are a few factors which should be taken into consideration as well, when interpreting the statistical agreement found herein. For example, the accuracy of the retrieved  $F_r$  and of  $T_s$  is a possible cause of error as the technique requires only those two parameters as inputs for its implementation. In our study, LST was measured by FLIR SC660 with an error lower than 0.03 °K, which is considered very small. Furthermore, since  $T_s$  is scaled in the “triangle”, the effect of the predicted temperature accuracy might be small (Carlson, 2007). Possible reasons for the lack of complete agreement could be related to the scale-mismatch between the EO-data and the in-situ measurements, geo-location errors, and surface heterogeneity at the UAV sensor spatial resolution, even though in this particular case predictions were obtained at very high spatial resolution. Another possible factor concerning the SSM comparisons in particular is that the ground measurements were acquired at 0 to 10 cm depth, while the UAV-derived ones respond to soil water content at a much shallower layer (0 to 5 cm) over bare soil. Effective soil depth for SSM measurement is an issue under investigation (Amani et al., 2016). Some studies (Finn et al., 2011; Kasim et al., 2020) suggest an effective measurement at a depth of 5 cm, while other studies (Zhang et al., 2015) suggest effective agreement at a depth of 10 cm. Furthermore, uncertainties due to the instrumentation accuracy for EF and  $R_n$  measurement should further be considered. Various studies have reported that errors in instantaneous LE flux measurement can be in the order of 20% to 30%, which can be even higher under certain circumstances (such as terrain features); similarly a measurement uncertainty for  $R_n$  of 10% is not uncommon (Petropoulos et al., 2013).

Despite the promising results obtained in this first investigation performed herein, the “simplified triangle” technique has some limitations which should also be acknowledged. Those include its requirement to have within the image field of view a sufficient variability of  $F_r$  and SSM range, in order to properly define the “wet” edge and the “dry” edge. Another issue is the possible human error in the selection of warm and cold edges. However, this is an issue common to other  $T_s$ /VI methods (Tomas et al., 2014; Mi et al., 2015). Furthermore, the technique assumes a linear relationship between the TS/VI feature space and the predicted EF and SSM, which might not necessarily be the case in nature.

Nonetheless, the “simplified triangle” capitalises on the inherent relationships existing in the  $T_s$ /VI feature space for estimating  $M_o$  and EF. Yet, it seems to have some strong advantages in comparison to other  $T_s$ /VI methods. The technique is simple to be applied and is dependent on a few input parameters which can be easily computed from EO sensors. This makes the technique implementation quick and computationally inexpensive when that is to be applied to small scale studies. Its implementation, particularly with UAV images, presents several advantages. When the technique is implemented with UAV data cloud cover is not an issue (as UAVs fly at very low altitude) as it would be if satellite data had been used. In addition, the technique when implemented with UAV data, the spatiotemporal variability of EF and  $M_o$  are computed at a very fine spatial resolution (at 4 cm in our case). As information on very high spatial and potentially temporal resolution of EF and SSM is essential to decision making in most agricultural applications, including precision agriculture (Wang et al., 2018; Cui et al., 2020), the potential added value of the “simplified” triangle technique to addressing this requirement is clear. In overall, all the above characteristics place the “simplified triangle” in a privileged position as a candidate for further investigation for a potential operationalisation with either with satellite or airborne EO data.

## 6. Conclusions

In this study, a first assessment of the so-called “simplified triangle” technique was performed to evaluate the ability of this method to predict EF and  $M_o/SSM$  when very high spatial resolution EO imagery acquired from UAV are available. A robust evaluation was carried out for an experimental site located in Sicily, Italy for which an extensive field campaign took place in July 2019. To our knowledge, the study represents the first detailed assessment of this innovative method with UAV data, particularly in a Mediterranean setting. The implementation of the investigated herein technique with UAV images presents several advantages. Data cloud cover is not an issue for UAV images and the spatiotemporal variability of EF and  $M_o/SSM$  are computed at a very fine spatial resolution (at 4 cm in our case). Regardless, UAV images present an additional challenge in correctly implementing the “simplified triangle” technique. The method requires a sufficient variability of  $F_r$  and SSM range within the image which can prove challenging in UAV imagery.

The obtained results suggest that the “simplified triangle” performed satisfactorily in predicting both the  $M_o/SSM$  and EF. Validation showed an average error of 0.053 for EF and of 0.040  $\text{cm}^3 \text{cm}^{-3}$  for SSM. This implies that the investigated approach performs well under the semi-arid conditions characterizing the experimental set up. Both predicted maps also exhibited sensible spatial variability across the experimental site, with patterns in agreement to land cover type, topography and other site-specific characteristics. The prediction accuracy of the technique was also in close agreement, or even better, than accuracies reported by other independent investigators using different  $T_s/VI$  approaches and EO data types.

However, the results reported herein are evidently based on a single image analysis. As the technique is recent, further scrutiny and additional studies are required to establish its applicability to different ecosystems. Such future investigations would require exploring the prediction accuracy of the technique in different ecosystem environments and for longer time periods using UAV imagery and spaceborne datasets from appropriate sensors (e.g. Landsat, Setinel 1 to 3, Moderate Resolution Imaging Spectroradiometer (MODIS)), as well as including a flux footprint analysis comparisons for the case of EF/ET predictions. In addition, a detailed sensitivity analysis of the method would also allow quantifying the effect of  $T_s$  and  $F_r$  errors on prediction accuracy. Other aspects of the technique that deserved investigation involve automating the process of determining the wet and dry edge, which would also eliminate user subjectivity in the technique implementation. It could potentially prove beneficial to combine pixels to satellite sensor spatial resolution (e.g. from the Landsat resolution of 30 or 120 m) to define the triangle boundaries. Then, once those boundaries have been established, they could be imposed on the higher resolution UAV image. All the above are topics of key importance that will be pursued in future studies.

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## 479 Author contributions

480 AM, GP, GC, and SM conceived and planned the experiments. AM and FC contributed to the  
481 experimental design of the spectroradiometric acquisitions and radiometric calibration of the images.  
482 GC, GP and SM coordinated the experiment and provided instrumentations. AM and GD  
483 contributed to the GNSS experimental design and processing. GP contributed to the experimental  
484 design and management of the soil moisture probes and processed the data. AM processed the flux  
485 tower data. SM and FC designed and acquired the UAV images. GPP, AP, TNC, DH and CC  
486 contributed to model implementation, results processing and analysis. GPP, AP prepared the  
487 original draft of the manuscript. All authors reviewed and edited the final version of the  
488 manuscript and contributed to the preparation of the revised manuscript.

489

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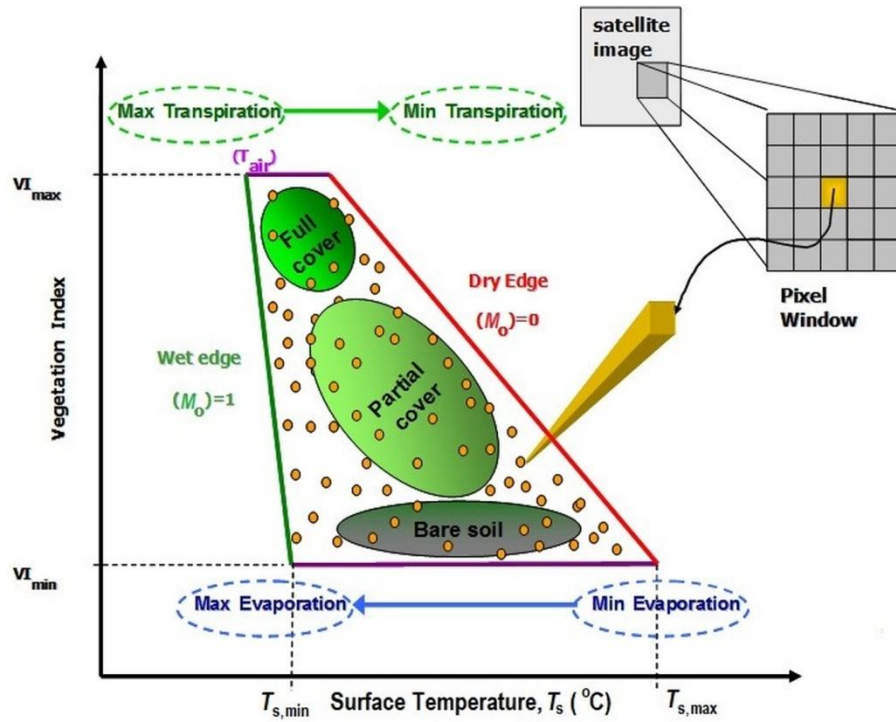
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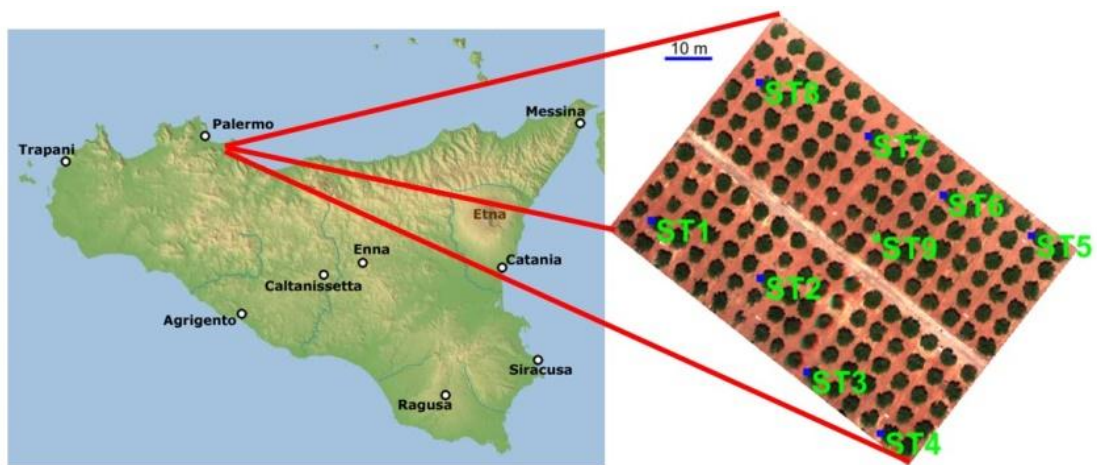
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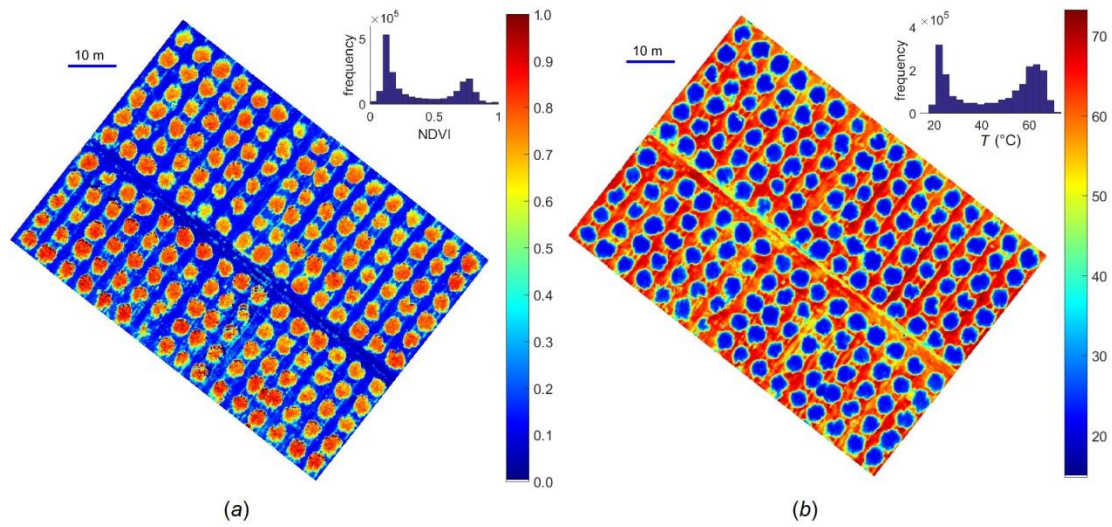




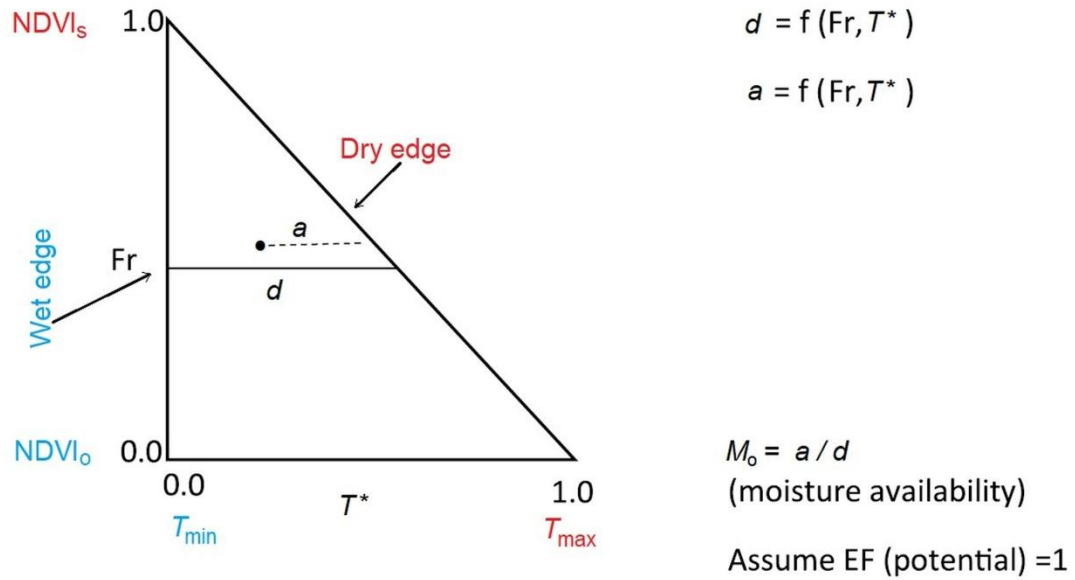
**Figure 1:** Conceptualisation of the main properties encapsulated in a  $T_s/VI$  scatterplot (adopted from Petropoulos et al. 2009).



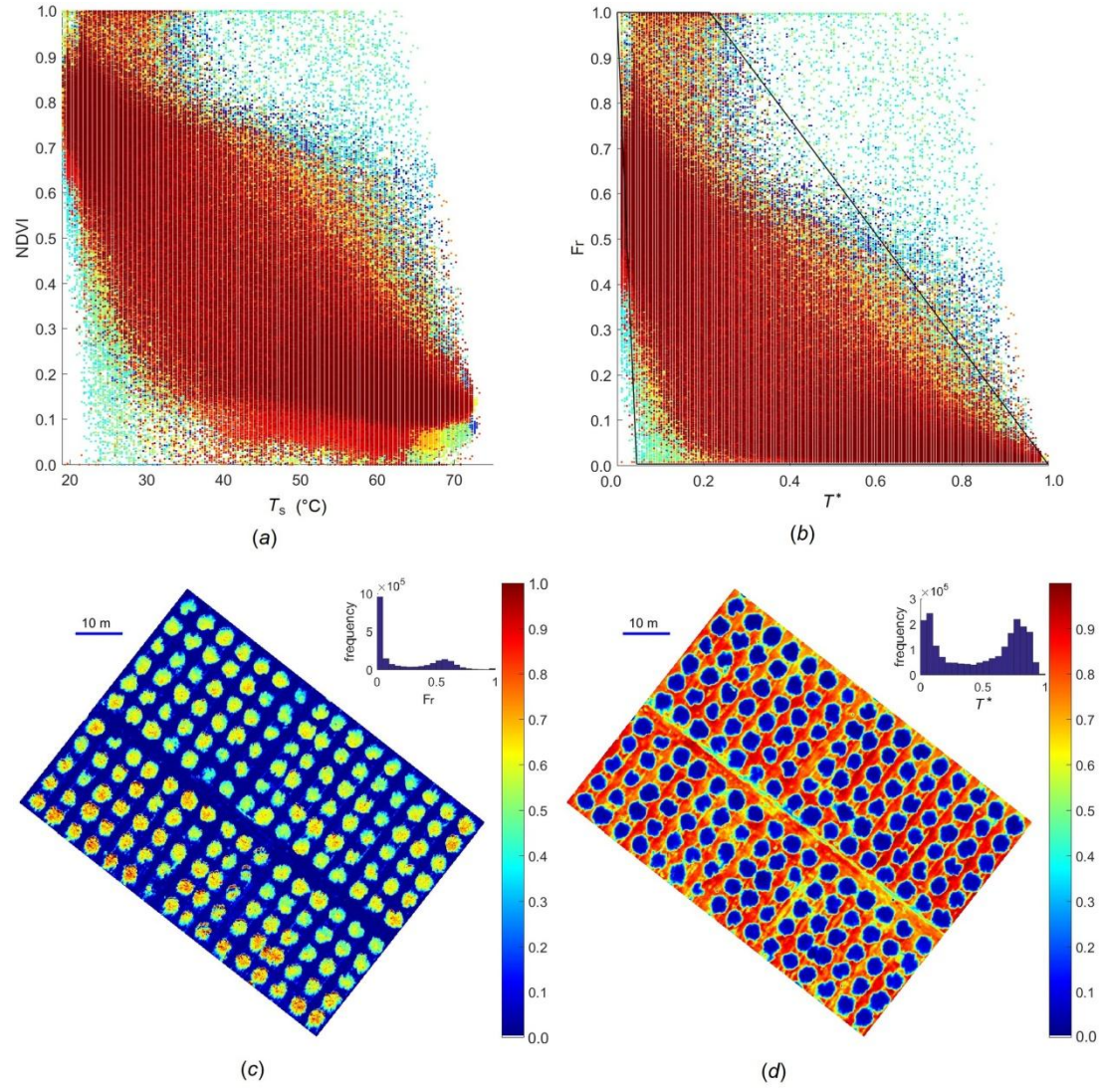
**Figure 2:** The experimental site, including the distribution of the ground measurement stations. ST1-8 refers to the locations of the probes that monitor soil water content, whereas ST9 is the eddy covariance system location. The image on the right is the actual UAV area covered by the UAV upon completion of orthorectification (see section 2.2.2 below).



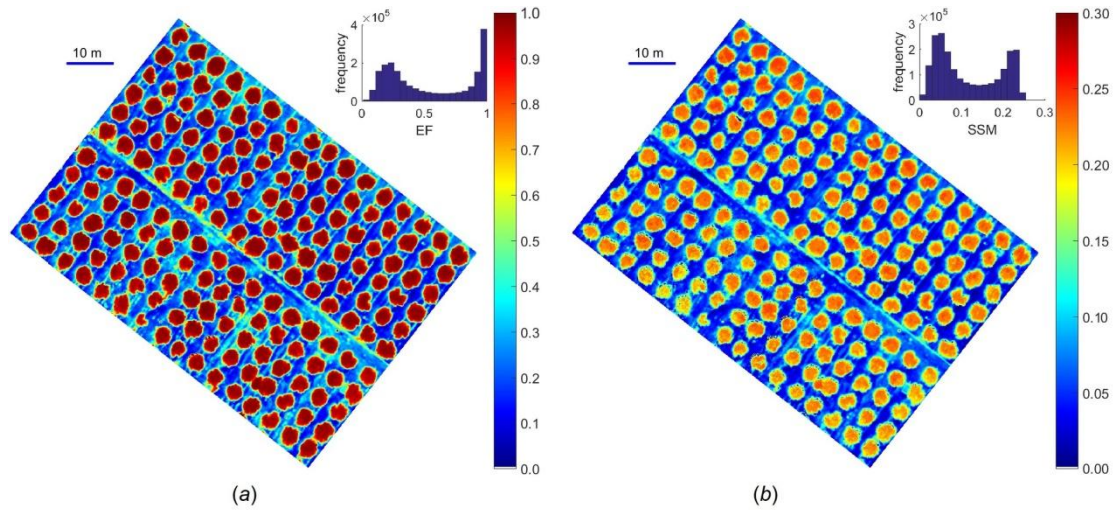
**Figure 3:** Pseudo colour maps of NDVI (a) and  $T_s$  (°C) (b) derived upon completion of the pre-processing steps. The insets show the frequency histograms of NDVI and  $T_s$  respectively. Temperature units are in Celsius.



**Figure 4:** Graphical summary of the “simplified” triangle method principles and critical points selection required in its implementation (adopted from Carlson & Petropoulos, 2019)

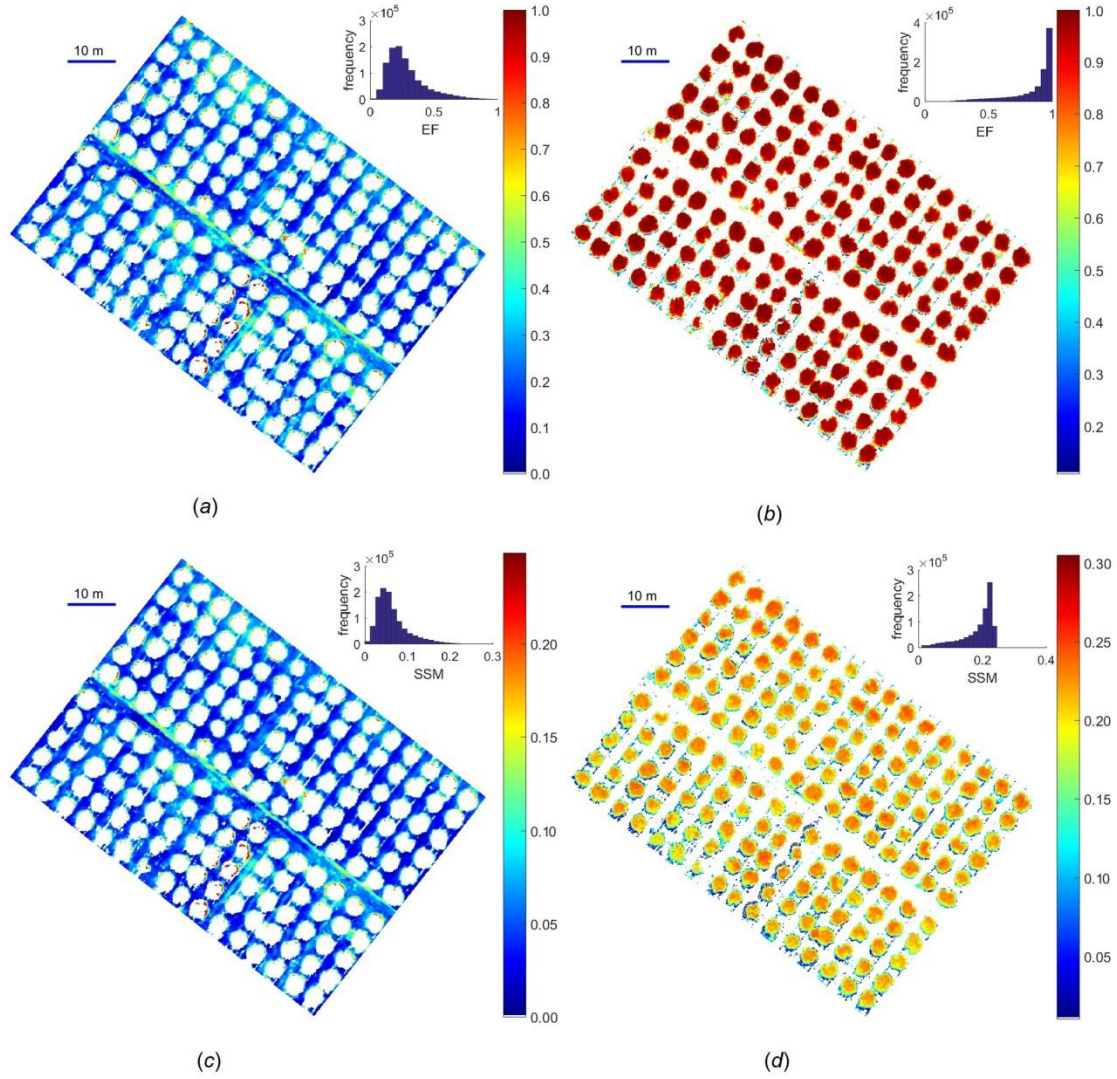


**Figure 5:** The scatterplots derived during the implementation (a,b), the  $F_r$  map (c) and the  $T^*$  map (d), derived from the datasets acquired with UAV. The “wet” and “dry” edge of the proposed triangle is shown by the continuous black line in scatterplot (b). The different colors in scatterplots (a,b) are for illustrative purposes only.

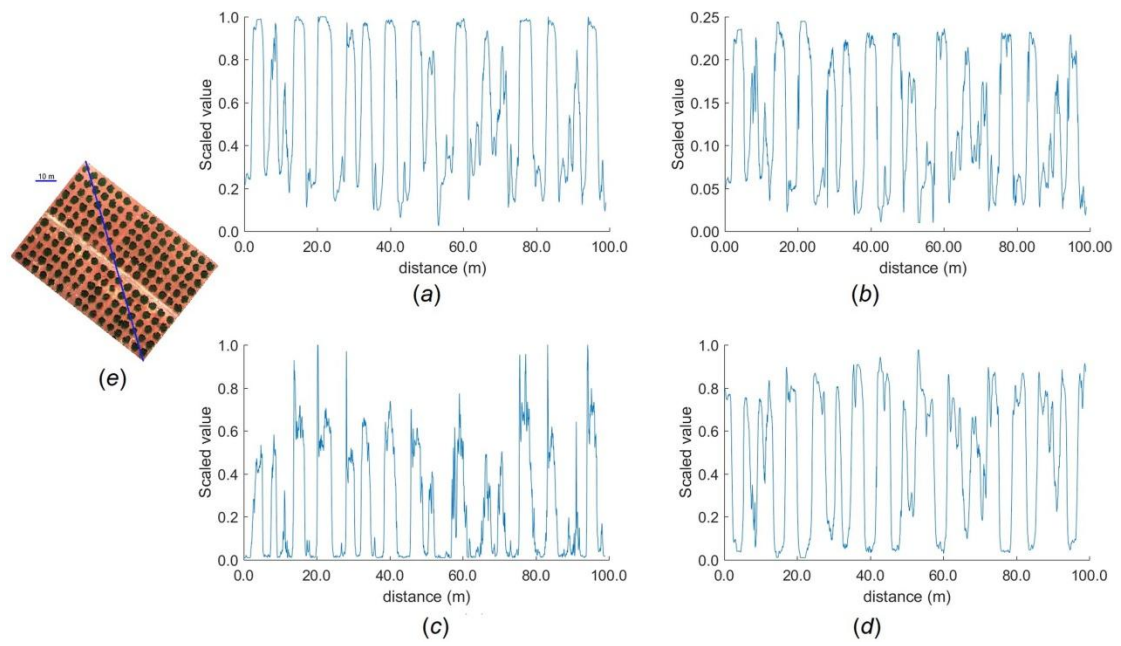




**Figure 6:** Maps of EF (a) and SSM (b) computed from the “simplified triangle” implementation using the data retrieved with UAV. The corresponding histograms are also shown.



**Figure 7:** EF maps computed separately for the vegetated area (b) and for bare soil (a). Similarly, the derived SSM maps for the vegetated area (d) and for the bare soil (c), are also shown. Each map is accompanied by the corresponding frequency histogram.



**Figure 8:** Arbitrarily selected transect within the field (e), and plots of the spatial variation of EF (a), SSM (b),  $F_r$  (c) and  $T^*$  (d) along the selected transect.

**Table 1:** Statistical measures used to assess the agreement between the predictions and ground observations. Subscripts  $i = 1 \dots N$  refer to the individual observations, while  $O$  and  $P$  refer to the observed and predicted values.

Name	Description	Mathematical definition
Bias / MBE	Bias (accuracy) or Mean Bias Error	$\text{Bias} = \sum_{i=1}^N \frac{P_i - O_i}{N}$
Scatter / SD	Scatter (precision) or Standard Deviation	$S = \sqrt{\sum_{i=1}^N \frac{(P_i - O_i - (\overline{P_i} - \overline{O_i}))^2}{N}}$
RMSE	Root Mean Square Error	$\text{RMSE} = \sqrt{\frac{\sum (P_i - O_i)^2}{N}}$

**Table 2:** Summary of the point by point comparisons between the ground observations ( $O$ ) and the corresponding predicted with the “simplified triangle” ( $P$ ). The differences ( $D$ ) between predicted and observed values are also indicated. Bias, Scatter and RMSE are expressed in units of  $\text{cm}^3 \text{cm}^{-3}$ .

Fluxes (-)	$O$	$P$	$D$
LE $R_n^{-1}$	0.266	0.319	0.053
SSM ( $\text{cm}^3 \text{cm}^{-3}$ )	$O$	$P$	$D$
SM1	0.139	0.090	-0.049
SM2	0.107	0.132	0.025
SM3	0.162	0.171	0.009
SM4	0.145	0.099	-0.045
SM5	0.078	0.073	-0.006
SM6	0.121	0.084	-0.037
SM7	0.145	0.084	-0.061
SM8	0.180	0.144	-0.036
		<b>Bias</b>	<b>-0.025</b>
		<b>Scatter</b>	<b>0.031</b>
		<b>RMSE</b>	<b>0.040</b>